# Thermohaline variability of the upper ocean: documenting the specific role of salinity in the Tropics



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## OUTLINE

General context

\* PART-1 On the impact of salinity on the stratification and

the dynamics of ENSO

the edge of the Western Pacific Warm Pool

the specific role of salinity stratification in N<sup>2</sup>

\*\* PART-2 On the importance of the small scale variability

emergence of interest

stirring by the mesoscale eddies: a case study

(relationships with the biogeochemical fields)

Conclusions

## From the large-scale and long-term variability of SSS...



But in situ Observations contain also some small scale variations (unsolved)





↑ indicates where climate change is predicted to increase the incidence or magnitude of that attribute,

 $\pmb{\Delta}$  indicates attributes where the impact of climate change on that attribute is  $\underline{variable}$ 



Examples will be shown for the

dynamics of ENSO & Western Pacific Warm Pool, detection of the edge of the WPWP,

the global variance of SSS

and salinity stratification in N<sup>2</sup>

#### Barrier layer thickness or Ocean salinity stratification



Definition of the BL : pycnocline controlled by the salinity stratification (halocline within quasi-isothermal layer)

## What are the effects of removing the BL in a coupled model framework?



Figure 2. Vertical profiles of temperature, salinity and density (left), and of the resulting profiles of Brunt-Väisälä frequency (right) when the salinity stratification is considered or not. Dashed horizontal lines indicate the upper and lower limits of the barrier layer.

We defined <u>pertubated experiments</u> (PERT) as :

From Maes et al. (2002)



(more details on the coupled model may be found in Belamari et al. *JClim* 2003)

N<sup>2</sup>(T,S) ≻ N<sup>2</sup>(T) : In the TKE vertical diffusion scheme when 4°N [SST > 28°C ] 4°S



What are the effects of removing the BL in a coupled framework?



Robust results obtained from the coupled model between the N<sup>2</sup>(T,S) and N<sup>2</sup>(T) simulations:

- 1. BLT plays a role on the depth of the mixed layer, with no systematic (cooling) impact on the SST (as may be expected),
- 2. BLT may affect the onset, the heat build-up and the mean state of the Pacific Ocean and ultimately ENSO dynamics
- 3. The most important O/A interactions for the warm pool region occur near its eastern edge (i.e., oceanic convergence zone),

# The edge of the WPWP

The ENSO «variability is generated primarily by a coupled ocean-atmosphere instability near the eastern edge of the western equatorial Pacific warm pool. » A. Clarke (2014)

Bosc et al. (2009) and Qu et al. (2014) confirm from direct observations based on Argo data that the BLT signal follows the expected dynamics near the WPWP.

BLT (in m).

2013

2012

2011

2010

2009

2008

2007

2006

140E

160E

180

160W

140W

2005 140W

SST (in °C). SSS 2013 -2013 Pt 1: SSS features are now well captured by satellite 2012 -2012products. 2011 2011 -Pt 2: Eastern edge of the 2010 -2010-WPWP is also marked in the ocean color and 2009 -2009 biogeochemistry fields. 2008 -2008 2007 2007 2006 2006 White lines mark the 29°C. 34.8 and 15 m values 2005 2005

140E

160E

180

160W

140W

140E

160E

180

160W

(Qu et al. 2014)

How well is the eastern edge of the WPWP reproduced by coupled models? The simulation of the WPWP edge is examined in 19 models from CMIP5



2. Results on detection of the edge of the WPWP



New metric – the warm pool edge is better simulated in CMIP5 than if assessed only on temperature. The eastern edge of the WPWP is crucial to represent the full dynamical response at the scale of the Pacific basin.

↑ indicates where climate change is predicted to increase the incidence or magnitude of that attribute,

 $\pmb{\Delta}$  indicates attributes where the impact of climate change on that attribute is variable



Examples will be shown for the

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#### CMIP5 ensemble mean fractional inband variance of internal SSS variability

SSS



O'Kane et al. JGR 2016



SST

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Monselesan et al. GRL 2016

2.3 Examples for the salinity stratification in N<sup>2</sup>

Q3: Are we able to determine the precise role of the salinity stratification?

Methodology:

$$N^{2}(T,S) = -\frac{g}{\rho} \frac{\partial \rho}{\partial z} \approx \left( g \alpha \frac{\partial T}{\partial z} - g \beta \frac{\partial S}{\partial z} \right)$$
$$= N^{2} \tau(T,S) + N^{2} s(T,S)$$

where

$$N^{2}_{T}(T,S) = g\alpha \frac{\partial T}{\partial z}$$

$$N^{2}_{S}(T,S) = N^{2}(T,S) - N^{2}_{T}(T,S)$$

$$\alpha = -\rho^{-1} \frac{\partial \rho}{\partial T}$$

$$\beta = \rho^{-1} \frac{\partial \rho}{\partial S}$$

We define the Ocean Salinity Stratification (OSS) as:

 $OSS = \langle N^2(T, S) - N^2(T) \rangle_{0-300m}$ where  $N^2(T, S) - N^2(T) > 0$  Q3: Are we able to determine the precise role of the salinity stratification?





#### PART-2: On the importance of the small scale variability



PART-2: On the importance of the small scale variability

What do we know about such variability in salinity ?

Tangible progresses have been achieved with recent satellite observations (SMOS & Aquarius), BUT...

 Microwave radiometers measure the salinity in the top few centimeters of the ocean

Satellites measure salinity as a spatial average over an area of about
 50 or 100x100 km<sup>2</sup>

- High resolution *in situ* data across large distance across the oceans are only available from TSG mounted on research vessels and ships.



## One example of the impact of eddies stirring: a study case in the Coral Sea



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![](_page_21_Figure_1.jpeg)

FSLE from the GEKCO surface currents (Sudre et al. 2013)

One example of the impact of eddies stirring /submesoscale structures

![](_page_22_Figure_1.jpeg)

## How to measure the impact of mesoscale eddies?

![](_page_23_Figure_1.jpeg)

### What is the role (in subsurface) of non linear eddies?

![](_page_24_Figure_1.jpeg)

## CONCLUSIONS

\* Salinity is seen as a proxy (indicator) for identifying climate-driven changes in global Earth's hydrological cycle (acceleration in response to climate change), BUT...

\* Salinity should be also viewed and considered as one (main) actor in the ocean response to climate forcing,

\* Ocean Salinity Stratification-OSS (stabilizing part of N<sup>2</sup> that is due to salinity only, i.e. also known as the barrier layer in some regions) is a natural variable of ocean dynamics (mean state, variability and ENSO dynamics...),

\* OSS can be significant in regions where both T and S are mixed over the same depth, and overall, the contribution of the OSS to N<sup>2</sup> could be as large as 40-50% of the contribution of the thermal stratification at the seasonal timescales,

\* Observation of mesoscale variability: eddies, turbulence, and lateral fluxes are improving, but it remains necessary to understand the growth and decay processes of the associated eddy structures that can also depend on sub-mesoscale dynamics.

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![](_page_27_Figure_0.jpeg)

SSS as a pertinent tracer of biogeochemical provinces: One example in the equatorial Pacific Ocean

Relationship with chl-a is also pertinent at the eastern edge of the Warm Pool (see also Maes et al., SOLA 2010)

OUTPACE cruise (18 Feb-03 Apr 2015) Southwest Pacific Ocean T. Moutin and S. Bonnet (PIs)

![](_page_28_Figure_1.jpeg)